

1 **Factors controlling recent diatom assemblages across a steep local nutrient gradient in central-eastern Finland**

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8

9 **Abstract**

10

11 The Iisalmi Route, a chain of lakes in central-eastern Finland, is more eutrophic than its surroundings. We used
12 multivariate analyses to study the influence of selected environmental factors (water quality, basin characteristics, and
13 subcatchment surficial geology) on the recent diatom assemblages across this steep local nutrient gradient. In addition
14 to the spatial analysis of surface sediment diatom assemblages from 51 sampling sites (48 lakes), temporal changes in
15 the total phosphorus (TP) concentrations of one Iisalmi Route lake (Lake Kirmanjärvi) were analyzed using weighted
16 averaging partial least squares regression and 27 fossil diatom samples. Both TP and electrical conductivity (EC)
17 showed statistically significant independent signals in the modern diatom data. The TP gradient was related to till grain
18 size variation suggesting that geological factors affect the spatial TP variation directly or indirectly through differences
19 in land use. Based on the temporal study, the direct effect of geology is most likely behind the steep nutrient gradient in
20 the area because Lake Kirmanjärvi was found to be naturally eutrophic. EC, on the other hand, seems related to
21 anthropogenic disturbance. Our study highlights the importance of taking the local geology into account when assessing
22 past or present water quality or planning for lake management.

23

24 **Keywords**

25

26 Lakes, eutrophication, phosphorus, electrical conductivity, Quaternary deposits, multivariate analysis, transfer function

27

28 **Introduction**

29

30 Eutrophication remains one of the most persistent problems affecting lake water quality after a nearly half a century of
31 mitigation practices (Ekholm & Mitikka, 2006; Schindler, 2006). Sediment-based paleolimnological techniques are
32 widely used for assessing human-induced changes as well as for estimating the pre-disturbance conditions of lakes in
33 order to set realistic management goals (Bennion et al., 2011a; Davidson & Jeppesen, 2013). Paleoecological methods
34 utilize remains of micro-organisms preserved in sediment cores. These data are commonly studied with multivariate
35 methods, suitable for revealing distribution patterns and the relationship between the species and environmental data,
36 and transfer functions that enable the quantitative reconstruction of past water quality from the paleoecological data
37 (Smol, 2008; Birks, 2010; Birks et al., 2012). Diatoms, in particular, are excellent biological indicators of
38 environmental conditions (Battarbee et al., 2001).

39

40 Diatom assemblages are controlled by a multitude of environmental factors with complex relationships (Juggins, 2013;
41 Rühland et al., 2015). These relationships tend to vary regionally as many characteristics of an individual lake system
42 are defined by the local geological and land use histories as well as climate-related factors, such as vegetation, thermal
43 properties, and insolation (Håkansson, 2005; Leavitt et al., 2009). The geological controls are based on basin
44 morphology and watershed properties, such as topography, soil erodibility, and nutrient availability. For example, in
45 regions where surficial geology is dominated by fine grained sediments, lake basins are typically shallow and the fertile,
46 easily eroded watersheds have attracted agriculture early in their history (Sherratt, 1980; Taavitsainen et al., 1998).

47

48 The Iisalmi Route, a mosaic of interconnected lakes in central-eastern Finland (Fig. 1), is a prime example of a region
49 where the relationship between environmental factors seems to differ from those in its surrounding areas. It has proven
50 to be a challenging locality for water management, because models that seek to predict water quality and lake response
51 to nutrient loading have performance issues, such as the underestimation of natural TP export (Dubrovin et al., 2016),
52 when applied to the area. The Iisalmi Route has been suspected to be naturally nutrient rich and, indeed, its flanks were
53 the first areas to be inhabited when central-eastern Finland was permanently populated in the 16th century (Soininen,
54 1961). Currently, it is characterized by conspicuously eutrophic lakes with generally poorer water quality than the lakes
55 in its vicinity (Kauppila et al., 2012).

56

57 The main purpose of this study was to examine the diatom species gradient and the spatial relationship between surface
58 sediment (recent) diatom assemblages and the environmental factors in the eutrophic Iisalmi Route and its nutrient
59 poorer adjacent areas with ordination and regression methods. In addition to water quality and basin characteristics, we
60 studied the influence of local geology (the composition of surficial sediments in the subcatchments of the lakes) to
61 diatom assemblages. Our aims were to find the environmental variables with the strongest statistical total and
62 independent effects on the diatom data, in other words the main drivers of the spatial diatom assemblage variation, and
63 to study this spatial variation in the diatom, water quality, basin, and catchment geology characteristics between the
64 Iisalmi route and its surrounding areas. Furthermore, we applied the found statistical connection between the modern
65 diatom assemblages and total phosphorus (TP) as a tool to study the past TP conditions of one Iisalmi Route lake (Lake
66 Kirmanjärvi) based on a previously published fossil diatom data set (Kauppila et al., 2012).

67

68 **Materials and methods**

69

70 *Lake selection and study area*

71

72 Lake selection for this study was done by first listing the lakes in the eutrophic Iisalmi Route and its nutrient poorer
73 surroundings in central-eastern Finland according to their water quality and then qualitatively selecting lakes so that the
74 different concentrations of the TP gradient from oligotrophic to hypertrophic were represented and that the variety of
75 catchment geologies was captured as well. The selection process was restricted by the availability of suitable water
76 quality observations, because the study area is remote and the water quality of many lakes is rather infrequently
77 monitored. As many sites were selected as was possible with the available resources. Small and shallow lakes are
78 typical in the area-of-interest (the Iisalmi Route) and, therefore, larger and deeper lakes were avoided in the lake
79 selection process. Also low pH lakes in swampy areas and lakes next to anthropogenic point sources were left outside of
80 this study.

81

82 The selected 48 lakes (51 sampling sites) are located between latitudes 62° 44' N-63° 52' N and longitudes 26° 18' E-
83 28° 96 04' E (Fig. 1). Most sampling sites (26) belong to the Iisalmi Route, 12 to the Rautalampi Route, 8 to the Nilsjä
84 Route, and 5 to the Kallavesi-Sorsavesi area. The Iisalmi Route is further divided into three subroutes (Kiuruvesi,
85 Salahmi, and Sonkajärvi) that all drain into a central basin area. The Iisalmi and Nilsjä Routes drain into the Kallavesi-

86 Sorsavesi area, all located in the upper reaches of the Vuoksi drainage system that flows into Lake Ladoga in Russia.
87 The Rautalampi Route, a headwater part of the Kymijoki drainage system, drains into the Gulf of Finland. The lakes are
88 mostly small (surface area < 10 km²) and shallow (average depth < 4 m) and the catchments vary in size from 0.1 to
89 5 600 km² (mean 605 km²) and subcatchments from 0.1 to 336 km² (mean 39 km²). The lakes are located at altitudes
90 ranging 82-160 m.a.s.l. and their retention time varies from 4.5 days to 3.5 years (mean 221 days). The Iisalmi route, in
91 particular, is characterized by small and shallow lakes that have high nutrient and humus concentrations (Table 1).

92
93 The bedrock in the area is characterized by the sheared Savo Belt that is located between the Archean Iisalmi complex
94 and the Proterozoic central Finland granitoid complex (Vaasjoki et al., 2005). These are overlain by Quaternary tills,
95 clays and other fine-grained sediments, eskers, and peat deposits. The tills covering Finland are typically coarse-
96 grained, but fine-grained till is also common in the study area. The area belongs to the southern boreal vegetation zone
97 (Ahti et al., 1968), but well-humified brown soils of the hemiboreal zone also exist in clayey and silty areas
98 (Taavitsainen et al., 1998). The annual mean temperature, precipitation, and evaporation are 3.2 °C (Pirinen et al.,
99 2012), 655 mm, and 464 mm, respectively (Korhonen & Haavanlammi, 2012).

100
101 The heaviest diffuse nutrient loading, mainly from grass cultivation, livestock husbandry, and forestry, is concentrated
102 on the Iisalmi Route and the downstream parts of the Nilsjä Route (Vallinkoski et al., 2016). Livestock breeding, in
103 particular, can lead to high amounts of easily available phosphorus (Järvenranta et al., 2014). In addition, many lakes
104 have been lowered artificially and the water levels of the larger lakes are regulated (Tanskanen, 2002; Vallinkoski et al.,
105 2016).

106
107 *Sampling and data sets*

108
109 Lake sediments were cored in late winter 2010 with a Limnos® gravity corer (Kansanen et al., 1991) that is designed
110 for taking undisturbed surface sediments. The coring sites were located in the deepest parts of the lake basins, at or near
111 pre-selected water quality monitoring sites of the environmental administration to provide a good match between water
112 quality data and sediment diatom assemblages. Diatom slides were prepared from the surface sediments (0-1 cm) by
113 using standard procedures described, for example, by Battarbee et al. (2001). A minimum of 400 valves were identified
114 with a phase-contrast light microscope at 1000x final magnification and by using the nomenclature of Krammer &

115 Lange-Bertalot (1986-1991), Krammer (1992), Lange-Bertalot & Moser (1994) and Houk (2003). The taxonomical
116 diversity of the diatom samples was calculated with Past software version 2.09 (Hammer et al., 2001) by using
117 individual rarefaction (Hurlbert, 1971). In addition to diatoms, chrysophyte cysts were counted from the slides and
118 transformed into cyst to diatom (CD) ratio (Werner & Smol, 2005).

119

120 Data for seven water quality variables and five lake basin characteristics were collected from the OIVA information
121 service and from VEMALA (Huttunen et al., 2016), the national-scale nutrient loading model for Finnish watersheds of
122 the Finnish Environment Institute (Table 1, see also Online Resource 1). When available, we calculated the medians of
123 the growing season (~May – mid-Oct.) water quality observations of 2000-2009. Small and remote lakes, typical to the
124 study area, are infrequently monitored, so 12 sampling sites lacked the desirable water quality data for one or more
125 environmental variable. For these lakes, we used either near growing season observations (Apr./late Oct.) of 2000-2009
126 or growing season observations from the 1990s. In three cases, a mixture of whole year observations from the two
127 decades had to be used. As TP was used for environmental reconstructions, we further ensured the applicability of the
128 used TP observations by drawing plots from all the available observations in order to see possible temporal trends. Most
129 lakes did not show any notable trend in TP during the last three decades and the variation of 2000-2009, in particular,
130 was mainly dominated by seasonal variation. In addition to the water quality and basin characteristic data, the
131 percentage of fields in each subcatchment was calculated based on the data available in VEMALA.

132

133 The distribution of surficial geology (Quaternary deposits) in the subcatchments of the study lakes (Table 2, see also
134 Online Resource 2) were calculated using ArcGIS software and the Quaternary maps of Finland (Geological Survey of
135 Finland, 2012). The most detailed (1:20 000) maps, covering approximately 50 % of the whole study area, were used
136 when available. For areas lacking the detailed maps, the 1:1 000 000 map was used for calculating the percentages of
137 fine-grained and coarse-grained till in the subcatchments and the 1:200 000 map (lacking a differentiation between the
138 two till types) for the other deposits.

139

140 For the temporal study we used a previously published fossil diatom data set from Lake Kirmanjärvi (Kauppila et al.,
141 2012) located in the Iisalmi Route (Fig. 1). The data set includes 27 core samples taken in 2006 with a downwards
142 increasing interval reaching down to 150 cm in the sediment. Based on the age-depth model of Kauppila et al. (2012),
143 the bottommost sample represents the mid-16th century at the latest.

144

145 *Numerical methods*

146

147 Prior to the ordination and regression analyses, the environmental data (excluding pH) were log-transformed to reduce
148 skewness in the original distributions and the species data were square root transformed to stabilize the variance of the
149 data (ter Braak & Šmilauer, 2002). Exploratory data analysis revealed one outlier sample (Lake Suo-Valkeinen) that
150 differed from the other samples based on its water quality, low pH (5.5) in particular, as well as its benthic diatom
151 community dominated by *Eunotia* spp., *Pinnularia* spp., and *Frustulia* spp, and a high CD ratio. Therefore, Lake Suo-
152 Valkeinen was excluded from further analysis.

153

154 The structure of the diatom data, as well as the relationship between the species and the environmental data (water
155 quality, basin characteristics, and Quaternary deposits), were examined with multivariate ordination methods by using
156 Canoco 4.5 for Windows software (ter Braak & Šmilauer, 2002). Detrended Correspondence Analysis (DCA; Hill &
157 Gauch, 1980) was first performed to examine the length of the species gradient and the properties of the unconstrained
158 axes. The length of the species gradient exceeded 1.5 SD units allowing the use of unimodal methods (ter Braak &
159 Prentice, 1988). Therefore, Canonical Correspondence Analysis (CCA; ter Braak, 1986) was used with Monte Carlo
160 permutation test (999 permutations) to find out how well the supplied environmental data explained the observed
161 variation in the species data and to discover the environmental factors with the strongest effects on the diatom data.
162 CCA was run separately for water quality and Quaternary deposit data. The relationships between the diatom data and
163 the water quality variables that had the strongest effects were further examined with partial, constrained CCA. In
164 addition, the species gradient lengths in the direction of the different environmental factors were determined by using
165 Detrended Canonical Correspondence Analysis (DCCA; ter Braak, 1986).

166

167 To further study the temporal changes in the nutrient concentrations of Lake Kirmanjärvi, we developed a diatom-TP
168 transfer function based on weighted averaging partial least squares regression (WA-PLS; ter Braak & Juggins, 1993)
169 and leave-one-out cross-validation (jack-knifing) with C2 software version 1.6.8 (Juggins, 2007). TP was chosen
170 because it had a stronger independent signal in the diatom data than TN. Diatom taxa present in at least three samples
171 and with a relative abundance of at least one percent, 110 in total, were included in the transfer function. The
172 reconstructed values for the past water quality conditions in Lake Kirmanjärvi were compared to the annual averages of

173 monitored values and LOESS smoothed annual averages, as well as to the sample scores of the first and second axes in
174 Lake Kirmanjärvi DCA. The fossil samples were also plotted passively on top of a CCA ordination formed by the
175 modern diatom and water quality data to see the trajectory of the lake during the last centuries in relation to modern
176 water quality in the area.

177

178 Furthermore, the fossil samples were tested for good modern analogs in the modern data set by calculating the squared
179 chord distances (SCD) between the fossil samples and their closest modern samples (Overpeck et al., 1985; Birks, 1995;
180 Simpson, 2012) with Python programming language. These minimum SCDs between each fossil sample and the
181 modern samples were compared to the 2.5th, 5th, and 10th percentiles of the distribution of the pairwise dissimilarities of
182 the modern samples (Bennion et al., 2004; Bennion et al., 2011a, 2011b; Simpson, 2012). Fossil samples with a
183 minimum SCD lower than the 2.5th percentile were considered to have very good modern analogs, whereas the 5th
184 percentile was used as a critical value for good modern analogs (Simpson, 2012). Fossil samples with a minimum SCD
185 exceeding the 10th percentile were regarded as having no modern analogs.

186

187 **Results**

188

189 *Diatom assemblages*

190

191 The diatom assemblages varied across the study area reflecting the variation in the environmental conditions between
192 the lake routes. The central basin area of the Iisalmi Route and the Kiuruvesi Subroute showed the least diverse diatom
193 communities with a high relative abundance of planktonic species and the lowest CD ratios (Fig. 2). Conversely, the
194 Kallavesi-Sorsavesi area was characterized by the most diverse diatom assemblages with equally abundant benthic and
195 planktonic taxa and a higher relative abundance of chrysophyte cysts than in the other areas.

196

197 From over 500 identified diatom taxa, the planktonic *Aulacoseira ambigua* (Grunow) Simonsen and *Aulacoseira*
198 *subarctica* (Müller) Haworth were the most abundant, particularly in the Iisalmi route. Small *Fragilaria* spp., such as *F.*
199 *construens* f. *venter* (Ehrenberg) Hustedt, *F. brevistriata* Grunow, and *F. exigua* Grunow, were the most abundant
200 benthic taxa. Other common species occurring in the whole study area included *Cyclotella stelligera* Cleve & Grunow,
201 *Asterionella formosa* Hassal, and *Aulacoseira lirata* (Ehrenberg) Ross. The areas surrounding the Iisalmi Route had

202 also relatively high abundances of *Tabellaria flocculosa* (Roth) Kützing, *Achnantes minutissima* Kützing, *Aulacoseira*
203 *distans* (Ehrenberg) Simonsen, *Aulacoseira distans* var. *nivalis* (Smith) Haworth, *Aulacoseira alpigena* (Grunow)
204 Krammer, *Cyclotella radiosae* Grunow, and *Cyclotella rossii* (Grunow) Håkansson. *C. rossii*, however, was only found
205 in the Rautalampi Route and in the Kallavesi-Sorsavesi area.

206

207 Less abundant *Aulacoseira granulata* (Ehrenberg) Simonsen, *Aulacoseira islandica* (Müller) Simonsen, and *Fragilaria*
208 *capucina* var. *gracilis* (Oestrup) Hustedt occurred primarily in the Iisalmi and Nilsjä Routes. *Cyclotella pseudostelligera*
209 Hustedt favored only a few lakes in the Kiuruvesi Subroute and in the Rautalampi and Nilsjä Routes, whereas *Diatoma*
210 *tenuis* Agardh was abundant only in Lake Kirmanjärvi and *Fragilaria nanana* Lange-Bertalot in Lake Saarisjärvi in the
211 Iisalmi Route.

212

213 *Ordination*

214

215 Unimodal ordination methods were used as the species gradient length (2.32 SD units) of the square root transformed
216 surface sediment diatom data (Lake Suo-Valkeinen excluded) was sufficient for their use. The results of the DCA based
217 only on the surface sediment diatom data and the CCA constrained to all water quality and basin characteristic variables
218 (Fig. 3) were very similar indicating that the selected variables were important in explaining the diatom data (Birks,
219 2010). The sum of all unconstrained DCA eigenvalues was 4.15, and the eigenvalues for the first and second axis were
220 0.26 and 0.16 explaining 6.3 % and 3.9 % of the variance in the species data, respectively. Correspondingly in the CCA
221 constrained to all water quality and basin characteristics, the eigenvalues for the first and second axis were 0.25 and
222 0.16 explaining 5.9 % and 3.8 % of the variation in the diatom data. The species-environment correlations of the first
223 and second CCA axes were 0.98 and 0.96. The CCA constrained to the Quaternary deposit data showed somewhat
224 lower eigenvalues (0.21 and 0.11), species-environmental correlations (0.94 and 0.92), and variances explained (5.2 %
225 and 2.7 %) for the first two axes than the water quality and basin characteristic data (Fig. 4). The sum of all canonical
226 eigenvalues was also higher for the water quality and basin characteristic data (1.10) than for the Quaternary deposits
227 (0.74).

228

229 The total effects of each environmental variable on the diatom data were studied with CCAs constrained to one
230 environmental variable at a time combined with a Monte Carlo significance test of the first canonical axis (marginal

231 testing). The relative explanatory power of a variable is often expressed as a λ_1/λ_2 ratio that is the ratio between the
232 eigenvalues of the first (λ_1 , constrained) and second (λ_2 , unconstrained) ordination axes (Juggins, 2013). The marginal
233 CCAs showed that all water quality factors, lake altitude, clay, fine- and coarse-grained tills, and other fine-grained
234 sediments ($p = 0.001$), in addition to the other basin characteristics ($p < 0.02$), were statistically significant ($p < 0.05$).
235 From these environmental variables, electrical conductivity (EC, $\lambda_1/\lambda_2 = 1.05$), total nitrogen (TN, $\lambda_1/\lambda_2 = 1.01$), lake
236 altitude from sea surface ($\lambda_1/\lambda_2 = 0.99$), and total phosphorus (TP, $\lambda_1/\lambda_2 = 0.96$) had the highest explanatory power.
237 Their species-environment correlation and explained variance in the species data were 0.94 and 5.1 % (EC), 0.93 and
238 4.8 % (TN), 0.90 and 4.4 % (altitude), and 0.95 and 4.9 % (TP). According to the DCCA, the gradient lengths in the
239 direction of these four variables were 2.30 SD units (EC), 1.89 SD units (TN), 1.73 SD units (altitude), and 1.80 SD
240 units (TP). The λ_1/λ_2 ratios, species-environment correlations, explained variances, and DCCA gradient lengths of the
241 statistically significant Quaternary deposits ranged between 0.65-0.87, 0.86-0.88, 3.4-4.0 %, and 1.13-1.48 SD units,
242 respectively.

243

244 The conditional or independent effects of the four variables with the highest λ_1/λ_2 ratios were further examined with
245 partial CCAs and Monte Carlo testing. In the partial CCAs, each of these variables were tested by first taking the other
246 three variables into account as covariables. All four tests were statistically significant ($p < 0.05$) but the independent
247 effects of EC (71 % of its total effect) and TP (63 %) were the strongest ($p = 0.001$). The eigenvalue of the first axis was
248 0.134 for EC and 0.114 for TP, species-environment correlations were 0.95 and 0.97, and explanative percentages 3.6
249 % and 3.1 %, respectively.

250

251 *Transfer function and reconstructions*

252

253 The WA-PLS regression-based diatom-TP transfer function (with two components) had a coefficient of determination
254 (r^2_{jack}) of 0.72 and a root mean squared error of prediction (RMSEP_{jack}) of 0.191 log $\mu\text{g l}^{-1}$ (Online Resource 3). The
255 diatom-inferred TP (DI-TP) of Lake Kirmanjärvi ranged mainly between 40-60 $\mu\text{g l}^{-1}$ (Fig. 5a). Two samples
256 representing the early 20th century showed elevated values up to 80 $\mu\text{g l}^{-1}$, but these samples lacked good modern
257 analogs in the modern data set. The DCA sample scores of the first and second axis showed increased variation in the
258 diatom assemblages since the late 19th century (Fig. 5a). The DI-TP correlated with the sample scores of the first DCA
259 axis (0.72, $p < 0.001$), but not with the second ($p > 0.5$). The reconstructions of the last few decades also resembled the

260 monitored TP values rather well (Fig. 5b) despite the 16th-19th century samples having better modern analogs than the
261 more recent samples (Fig. 5a). The 2.5th, 5th, and 10th percentiles of the pairwise dissimilarity distribution of the modern
262 samples in the transfer function training set (used in the analogue matching) were 0.35, 0.39, and 0.47, respectively. The
263 trajectory of the fossil samples passively plotted on top of a CCA ordination of the modern data set displayed a recent
264 increase in EC and pH in contrast to the rather similar nutrient conditions at the core ends (Fig. 6).

265

266 Discussion

267

268 *Spatial variation in diatom assemblages and water quality*

269

270 The spatial variation in the diatom assemblages is notable within our study area despite the limited geographical
271 coverage, and the selected water quality variables are important in explaining this variance according to the ordination
272 analysis. The Iisalmi Route samples, those in the Kiuruvesi Subroute and the central basin area in particular, plot to the
273 high ends of the nutrient gradients in the CCA ordination (Fig. 3). These lakes are also humic, shallow, and have a low
274 Secchi depth and TN:TP ratio. In contrast, the Kallavesi-Sorsavesi area and the Rautalampi Route samples plot to the
275 opposite ends of these gradients. Correspondingly, the abundance of benthic diatom taxa, number of taxa, and CD ratio
276 are lowest in the Kiuruvesi Subroute (25 %, 63, and 9, respectively) and in the central lake area of the Iisalmi Route (15
277 %, 48, and 7), whereas they are highest in the Kallavesi-Sorsavesi area (50 %, 81, and 29). The abundance of planktonic
278 taxa in relation to benthic taxa, low species diversity, and a low CD ratio are often considered as indicators of eutrophy
279 (e.g. Smol, 1985; Wetzel, 2001; Vadeboncoeur et al., 2003).

280

281 The variation in the diatom assemblages along the TP gradient is similar to that in the southern and eastern Finnish
282 surface sediment data sets of Kauppila et al. (2002) and Miettinen (2003) that cover larger geographical areas than our
283 data set and represent, in general, deeper and less humic lakes with lower nutrient concentrations and higher EC. *A.*
284 *ambigua* and *A. subarctica*, that thrive in shallow, turbid, and eutrophic lakes (Gibson et al., 2003; Kangur et al., 2007;
285 Heinsalu et al., 2008), as well as small, benthic *Fragilaria*, common in shallow lakes (Bennion, 1994; Sayer, 2001), are
286 widespread and abundant in all three data sets. The upper ends of the TP gradients in these three data sets are
287 characterized by eutrophic taxa, such as *A. granulata*, *Cyclostephanos dubius* (Fricke) Round, *Fragilaria ulna* (Nitzsch)
288 Lange-Bertalot, and *Fragilaria crotonensis* Kitton (Van Dam et al., 1994; Bradshaw & Anderson, 2003; Bennion et al.,

289 2015; Ventelä et al., 2016). *A. ambigua*, *A. subarctica*, *A. granulata*, *C. dubius*, and *Fragilaria* spp., in particular, are
290 common taxa in the modern as well as background conditions of naturally eutrophic Finnish lakes (Räsänen et al.,
291 2006). In contrast, oligo-mesotrophic species, such as *C. rossii*, *A. distans* with its subspecies, *T. flocculosa*, *C. radiosa*,
292 and certain benthic taxa (e.g. *Brachysira* spp.) are common in the lower ends of the TP gradients in the three surface
293 sediment data sets. The most striking difference between the data sets is in the abundance of *Stephanodiscus hantzschii*
294 Grunow that is infrequent in our data set but typical in the high TP and EC lakes of fertile clayey or silty areas
295 (Miettinen, 2003), such as those in the southern Finnish data set of Kauppila et al. (2002).

296
297 According to the CCA variance partitioning, both TP and TN have a statistically significant independent effect on the
298 diatoms but the independent fraction of TP (63 %) from the variance explained by the total effect of TP is higher than
299 that of TN (58 %). The remaining fraction is shared with other confounding variables and can be either the result of a
300 direct ecological effect or the environmental variable-of-interest acts as a surrogate for another, correlated
301 environmental variable (Juggins et al., 2013). The λ_1/λ_2 value of TP is below 1.5 in our data set, which suggests that it
302 includes a mixture of taxa with responses to TP as well as to other confounding variables (Juggins, 2013). The λ_1/λ_2 of
303 TP is also slightly lower than the optimal value for reconstruction ($\lambda_1/\lambda_2 \geq 1.0$; ter Braak, 1988), but the similarity in
304 taxon responses along the TP gradient in our model and the two models that cover geographically larger areas suggest
305 that TP is an ecologically important factor affecting the diatom assemblages (Juggins et al., 2013).

306
307 Indeed, most diatom taxa in our spatial data set appear to follow the TP gradient, but certain cyclotelloid, elongated
308 pennate, and small *Fragilaria* taxa occur throughout the study area either in most lakes (e.g. *A. formosa*, *C. stelligera*,
309 and *F. construens* f. *venter*) or sporadically in a few lakes (e.g. *D. tenuis*, *F. crotonensis*, and *C. pseudostelligera*) and
310 are likely driven by factors other than TP. Small cyclotelloid and elongated pennate diatoms have been associated with
311 a reduction in ice cover and an increase in thermal stability related to anthropogenic climate warming (Rühland et al.,
312 2015), whereas small, benthic *Fragilaria* are common in shallow lakes and can be more sensitive to habitat availability
313 than to TP (Bennion, 1994; Sayer, 2001). Anthropogenic climate warming in Finland exceeds the global warming trend
314 and has led to earlier ice break-up, later freezing, and shorter ice cover duration in lakes (Korhonen, 2006; Tietäväinen
315 et al., 2010; Mikkonen et al., 2015). In Lake Pörovesi (a central basin of the Iisalmi Route), the period of permanent ice
316 cover has shortened approximately by a month since the late 19th century. However, many of our lakes are colored and
317 could, therefore, be less sensitive to climate warming than clear water lakes (Snucins & Gunn, 2000).

318

319 In addition to nutrients, EC has a statistically significant independent effect of 71 % on the diatom assemblages of our
320 spatial data set, but it does not distinguish the lake routes as clearly as TP. The upper end of the EC (and pH) gradient is
321 characterized by elongated pennate and small cyclotelloid taxa (e.g. *D. tenuis*, *C. dubius*, *C. pseudostelligera*, *F. ulna*,
322 and *F. crotonensis*), whereas *Aulacoseira* taxa are typical in the lower gradient end. *D. tenuis* and *F. crotonensis*, in
323 particular, are only or more commonly found in the recent sediments of naturally eutrophic Finnish lakes than in their
324 pre-disturbance conditions suggesting that these taxa are influenced by anthropogenic activities (Räsänen et al., 2006).
325 *D. tenuis*, for example, is most likely related to road de-icing in our study area. It is only abundant in the three lakes
326 with the highest EC (Lake Siilinjärvi, Lake Kirmanjärvi, and Lake Pöljänjärvi) that are located next to a main road and
327 groundwater areas with elevated chloride concentrations due to road salt (Vallinkoski et al., 2016). In addition, the
328 supplementary variable of subcatchment field percentage resembles the EC gradient despite the EC values in our data
329 set (mean 4.8 mS m⁻¹) being low compared to the typical Finnish values (5-10 mS m⁻¹) and especially to those increased
330 by the intensive fertilization of fields (15-20 mS m⁻¹; Oravainen, 1999). Therefore, agriculture and other anthropogenic
331 activities seem to influence the diatom assemblages in a way that is reflected in the EC gradient, whereas the TP
332 gradient appears to be controlled by the underlying geology.

333

334 *Spatial variation in Quaternary deposits*

335

336 The occurrence of many diatom taxa in our data set follow the boundaries of the lake routes suggesting that the steep,
337 local nutrient gradient is related to local geology either directly or indirectly through differences in land use (agriculture
338 vs. forestry). Furthermore, the CCA sample and species biplots of the Quaternary deposits (Fig. 4) and water quality
339 (Fig. 3) are strikingly similar. The subcatchments of the lakes in the eutrophic Iisalmi Route are characterized by a high
340 percentage of fine-grained sediments including fine-grained till, clay and other fine-grained sediments, such as gyttja
341 and silt. Due to these easily workable, fine-grained soils, the flanks of the Iisalmi Route were particularly attractive
342 already for the early settlers of eastern Finland (Soininen, 1961; Orrman, 1991; Taavitsainen et al., 1998). In contrast,
343 subcatchments in the least eutrophic Kallavesi-Sorsavesi area and the Rautalampi Route have high abundances of
344 coarse-grained till, bedrock outcrops, and peat deposits.

345

346 The spatial distribution of the fine-grained and coarse-grained basal tills in the study area is the result of glacial
347 dynamics during the retreat of the ice margin approximately 9 500 years ago. The fine-grained till, with a high clay
348 fraction up to 30 % (Kukkonen & Sahala, 1988), forms a wedge-shaped area stretching from the Bothnian Bay coast
349 across our study area in NW-SE direction. The fine fraction of the till originates from material eroded from interglacial
350 fluvial sediments deposited in the Gulf of Bothnia (Lintinen, 1995). The fine-grained till deposited under a passive ice
351 sheet in an interlobate area between two actively flowing, fan-shaped ice lobes, the North Karelian and the Lake District
352 ice lobes (Punkari, 1997; Lunkka et al., 2004). The coarse-grained till areas, on the other hand, formed under these
353 active ice lobes, where meltwater washed away the finest fractions from the till (Lintinen, 1995). Hence, there was net
354 erosion in the active ice lobe areas and net accumulation in the interlobate area of passive ice (Punkari, 1997).

355

356 The abundance of clay in the subcatchments has a notably similar signal in the diatom data than the passively plotted
357 subcatchment field percentage indicating that agriculture is still mostly located on the clayey flanks of the present-day
358 lakes. These clays deposited on the bottom of the previous stages of the Baltic Sea and large lakes (ancient Great Lake
359 Saimaa and Lake Päijänne) that formed after the Ancylus Lake drained. Kukkonen & Sahala (1988), for instance, found
360 from Lake Porovesi (a lake collecting the waters from the three subroutes of the Iisalmi route) a complete sediment
361 sequence since the deglaciation, spanning from Yoldia Sea varved clays and silts through Ancylus Lake clays and silts
362 and Great Lake Saimaa clays to the present-day Lake Porovesi clays. The other fine-grained deposits in the area are
363 most likely sediments of artificially lowered or dried lakes (Kukkonen & Sahala, 1988).

364

365 A comparison between the water quality and Quaternary deposit CCA plots (Fig. 3 & 4) also shows that the lakes with
366 the lowest pH are situated on subcatchments that have the highest abundances of peatlands. The three lakes with the
367 highest EC, on the other hand, are located next to an esker that formed between the active and passive ice lobes and
368 runs in a NW-SE direction along the eastern flanks of the Iisalmi Route central basins. These three relatively small
369 lakes appear to have been more susceptible to the effects of road salt from the de-icing of the main road than the larger
370 central basins. However, the three lakes do not plot to the high end of the sand/gravel gradient, most likely because the
371 esker with the main road is not the only sand and gravel formation in the area.

372

373 *Temporal variation in Lake Kirmanjärvi*

374

375 The fossil diatom assemblages, the DI-TP reconstructions, and the ordination analysis all indicate that Lake Kirmanjärvi
376 has been eutrophic prior to intensive human impact in the area suggesting that the steep nutrient gradient in the study
377 area is due to the direct effect of underlying geology. The mid-16th century core bottom sample most likely represents
378 conditions close to the natural state of Lake Kirmanjärvi as permanent population spread into the area in the early 16th
379 century (Soininen, 1961) and the first road was built in the late 18th century (Eskelinen, 1985). According to Kauppila et
380 al. (2012), there are no major changes in the diatom assemblages of Lake Kirmanjärvi prior to 1836 CE (local diatom
381 assemblage zone (LDZ) 3). The DCA sample scores (Fig. 5a) remain relatively steady until 1875 CE indicating a rather
382 stable diatom assemblage. Therefore, the pre-1875 (or at least pre-1836) conditions could be considered as the reference
383 conditions of Lake Kirmanjärvi. This is in accordance with the conclusion of Battarbee et al. (2011) that pre-1850
384 conditions are a suitable reference state for lake restoration in most European lakes. The pre-disturbance samples have
385 also closer modern analogs in our spatial data set than the more recent samples (Fig. 5a). Lake Onkivesi, the central
386 basin into which the Iisalmi Route drains, is the closest modern analog for most pre-1875 samples of Lake Kirmanjärvi,
387 whereas Lake Nerkoönjärvi, the central basin that receives waters from Lake Kirmanjärvi and drains into Lake
388 Onkivesi, is the closest modern analog for most post-1875 samples.

389
390 The pre-disturbance DI-TP of Lake Kirmanjärvi is high (average 52 $\mu\text{g l}^{-1}$), approximately 15-20 $\mu\text{g l}^{-1}$ higher than with
391 the southern Finnish transfer function of Kauppila et al. (2002) in which humic lakes, such as those in the Iisalmi Route,
392 were excluded. However, it fits the upper end of the background concentration range in the previously studied naturally
393 eutrophic southern Finnish lakes (median 35-40 $\mu\text{g l}^{-1}$; Räsänen et al., 2006) and is somewhat higher than the pre-
394 disturbance DI-TP of Lake Onkivesi (40 $\mu\text{g l}^{-1}$; Miettinen, 2005). The diatom taxa common throughout the Lake
395 Kirmanjärvi core as well as in the modern spatial data set, such as *A. ambigua*, *A. subarctica*, *A. islandica*, *A. granulata*,
396 and *C. dubius* (Kauppila et al., 2012), also correspond to the taxa abundant in the background conditions of naturally
397 nutrient rich Finnish lakes (Räsänen et al., 2006). The natural eutrophy of the lake is in accordance with the extremely
398 high percentage (84 %) of fine-grained sediments in its subcatchment that have also led to the relatively high pre-
399 disturbance EC and pH (Fig. 6).

400
401 The high correlation between the DI-TP record and the DCA sample scores of the first axis (Fig. 5a) suggests that TP
402 has been an important factor affecting the Lake Kirmanjärvi diatom assemblages during the past centuries. The DI-TP
403 ranges mostly between 40-60 $\mu\text{g l}^{-1}$ throughout the core, but there are elevated concentrations up to 70-80 $\mu\text{g l}^{-1}$ during

404 the early 20th century simultaneously with elevated fine-grained mineral matter input levels that reflect increased
405 erosion and sedimentation most likely related to modern agriculture and deforestation (Kauppila et al., 2012).
406 Population growth led to the crisis of slash-and-burn cultivation in the end of the 19th century and to the subsequent
407 transition to modern agriculture with mineral fertilizers, more effective ploughs, and the ditching of fields, forests, and
408 peatlands (Soininen, 1974). The DI-TP returns to its background level during the 1960-1970s, whereas the diatom
409 assemblage shows a shift from LDZ 2 into LDZ 1 during this time with a particularly notable increase in the abundance
410 of *F. crotonensis* as well as *C. stelligeroides* (Fig. 5; Kauppila et al., 2012). The increase in the abundance of small
411 cyclotelloids and elongated pennates suggests a possible link to anthropogenic climate warming as proposed by
412 Rühland et al. (2015), but the area is also under the impact of modern agriculture. The field percent in the Lake
413 Kirmanjärvi subcatchment is currently the highest of our spatial data set (33 %).

414
415 The temporal changes in Lake Kirmanjärvi support the interpretation that the relatively steady TP concentrations reflect
416 the local geology (apart from the elevated values of the early 20th century), whereas a recent increase in EC and pH
417 since the mid-1990s (Fig. 6, also seen in water quality monitoring) is most likely related to road salt and the location of
418 the lake next to the main road. This increase in EC and pH is simultaneous with a further increase in the abundance of
419 small cyclotelloids and elongated pennates, such as *D. tenuis* and *C. pseudostelligera* (Kauppila et al., 2012). The mid-
420 1990s water quality change, as well as the lack of good modern analogs for certain post-disturbance samples, could
421 confound the DI-TP reconstructions of the topmost samples, but the DI-TP values of the last few decades are of the
422 same order of magnitude as the monitored values (Fig. 5b). Nevertheless, it needs to be taken into account that diatoms
423 are more reliable indicators of long-term trends than short-term variation.

424

425 **Conclusions**

426

427 The conspicuously eutrophic Iisalmi Route in central-eastern Finland clearly differs from its surrounding areas based on
428 the multivariate analyses of recent diatom assemblages, lake water quality and basin characteristics, and the surficial
429 sediment distribution in the lake subcatchments. The Iisalmi Route is characterized by nutrient rich lakes, diatom taxa
430 typically associated with eutrophic lakes, and a high abundance of fine-grained sediments (mainly fine-grained till and
431 clay) in the subcatchments. Conversely, the least eutrophic neighboring catchments are located in the eastern and
432 southern parts of our study area where coarse-grained till, bedrock outcrops, and peat are common. The lakes in these

433 catchments have relatively low nutrient concentrations and the diatom assemblages are typical of oligo-mesotrophic
434 lakes. The distribution of fine and coarse-grained tills in the area is the result of ice sheet dynamics during the
435 deglaciation of the continental ice sheet c. 9 500 years ago.

436

437 The nutrient concentrations seem directly or indirectly affected by the variation in the local geology because the
438 occurrence of many diatom taxa follows the boundaries of the lake routes and the nutrient gradients follow the
439 distribution of fine-grained versus coarse-grained sediments. The indirect effect of geology is related to the better
440 suitability of fine-grained sediments for cultivation (agriculture vs. forestry). However, the temporal study of Lake
441 Kirmanjärvi indicates that the lake has been eutrophic already prior to intensive human impact suggesting that the
442 diatom assemblages are, indeed, driven by the direct effect of the local geology. On the other hand, some diatom taxa
443 occur throughout the spatial dataset suggesting that they could be influenced by other factors than the local geology and
444 nutrients. In particular, certain small cyclotelloid and elongated pennate taxa, indicative of high electrical conductivity,
445 pH, and nutrient conditions, seem to be related to anthropogenic actions, such as agriculture, road de-icing, or climate
446 warming. Also the shallowness of the lakes is visible in the diatom data as relatively high abundances of small, benthic
447 *Fragilaria* taxa.

448

449 This study highlights the importance of assessing the geology of an area, in addition to the typical examination of water
450 quality, basin characteristics, and aquatic ecology, when evaluating the present or past conditions of a lake or planning
451 for possible management measures. Our results indicate that the distribution of different Quaternary deposits in an area
452 with a geographically limited coverage can lead to a steep local nutrient gradient. Because temporal changes in the
453 fossil diatom assemblages and DI-TP were assessed in one lake, further studies are needed to verify whether the other
454 currently nutrient rich lakes in the Iisalmi Route are also naturally eutrophic.

455

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466

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711 **Tables**

712

713 **Table 1** Summary statistics (minimum, average, and maximum values) for the lake water quality and basin

714 characteristics in the whole study area and averages for the different lake routes

	Min	Av	Max	Av ^{ir}	Av ^{ks}	Av ^{sa}	Av ^{ss}	Av ^{ic}	Av ^{nr}	Av ^{ka}	Av ^{rr}
TP ($\mu\text{g l}^{-1}$)	3.0	46	120	67	81	93	41	59	24	16	27
TN ($\mu\text{g l}^{-1}$)	150	772	1750	980	1105	1160	756	922	574	464	583
TN/TP	9	24	88	16	14	13	19	16	31	42	30
pH	5.5	6.9	7.8	7.0	7.0	7.5	6.8	7.1	6.8	6.7	6.9
EC (mS m^{-1})	0.5	4.8	12	5.4	5.7	6.6	3.9	6.1	5.2	3.3	3.8
Color (mg Pt l^{-1})	5	117	340	156	185	58	144	146	73	59	87
Secchi depth (m)	0.4	1.5	8.0	0.9	0.7	1.1	1.1	0.9	2.3	2.0	2.1
Sampling depth (m)	1.2	8.2	25	6.5	5.1	3.4	10	6.0	8.2	15	9.2
Average depth (m)	0.7	2.8	6.9	2.3	2.0	1.1	2.9	2.8	3.5	4.0	2.8
Area (km^2)	0.039	9.0	110	13	6.6	0.78	2.9	49	4.0	0.56	6.2
Volume (10^6 m^3)	0.11	27	360	41	18	0.83	8.4	160	14	1.9	15
Altitude (m.a.s.l.)	82	109	160	99	101	86	101	96	109	138	120

715 TP = total phosphorus, TN = total nitrogen, EC = electrical conductivity

716 ^{ir} = Iisalmi Route, ^{ks} = Kiuruvesi Subroute, ^{sa} = Salahmi Subroute, ^{ss} = Sonkajärvi Subroute, ^{ic} = Iisalmi Route central

717 basin area, ^{nr} = Nilsjä Route, ^{ka} = Kallavesi-Sorsavesi area, ^{rr} = Rautalampi Route

718

719

720 **Table 2** Summary statistics (minimum, average, and maximum values) for the Quaternary deposits and the percentage
 721 of fields in the study lake subcatchments of the whole study area and averages for the different lake route areas

	Min	Av	Max	Av ^{ir}	Av ^{ks}	Av ^{sa}	Av ^{ss}	Av ^{ic}	Av ^{nr}	Av ^{ka}	Av ^{rr}
F-g till (%)	0	29	77	38	51	46	12	42	26	23	15
Clay (%)	0	9.7	43	17	15	37	18	12	6.0	0	0.07
Other fines (%)	0	6.1	45	6.9	4.1	8.2	1.5	20	11	2.1	2.7
Sand/gravel (%)	0	3.9	26	1.9	1.6	1.7	1.3	3.5	5.2	7.1	6.0
Bedrock (%)	0	9.7	48	8.3	9.5	5.1	9.9	4.3	9.0	13	12
C-g till (%)	0	29	84	18	9.4	0.3	46	6.0	22	42	51
Peat (%)	0	12	77	8.9	8.1	2.6	11	11	20	11	13
Fields (%)	0	11	33	14	15	24	8.3	15	11	3.1	5.8

722 F-g till = fine-grained till, C-g till = coarse-grained till

723 ^{ir} = Iisalmi Route, ^{ks} = Kiuruvesi Subroute, ^{sa} = Salahmi Subroute, ^{ss} = Sonkajärvi Subroute, ^{ic} = Iisalmi Route central
 724 basin area, ^{nr} = Nilsjä Route, ^{ka} = Kallavesi-Sorsavesi area, ^{rr} = Rautalampi Route

725

726 **Figure captions**

727

728 **Fig. 1** Map of the study area in central-eastern Finland showing sampling sites, lake route catchments (thick lines), the
 729 subroute catchments of the Iisalmi Route (moderately thick lines), the subcatchments of individual lakes (thin lines), the
 730 drainage direction of the lake routes, agricultural land, and the location of Lake Kirmanjärvi (empty box) whose
 731 catchment area is shown in more detail in the inset

732

733 **Fig. 2** Relative abundances of selected diatom taxa, rarefaction, and chrysophyte cyst to diatom (CD) ratio in 51 surface
 734 sediment samples grouped according to the lake routes

735

736 **Fig. 3** CCA biplot of **a**) water quality variables and basin characteristics (arrows) with sampling sites in Kiuruvesi
 737 Subroute (squares), Salahmi Subroute (triangles), Sonkajärvi Subroute (filled diamonds), Iisalmi Route central basins
 738 (empty diamonds), Nilsjä Route (filled stars), Kallavesi-Sorsavesi area (empty stars), and Rautalampi Route (black
 739 circles) and **b**) diatom taxa (gray circles) and the environmental variables. The percentage of fields in a subcatchment
 740 (dashed arrow) is plotted passively on top of the ordinations. See Online Resource 1 for further information on the
 741 sampling sites and their water quality

742

743 **Fig. 4** CCA biplot of **a**) Quaternary deposits (arrows) and sampling sites in Kiuruvesi Subroute (squares), Salahmi
744 Subroute (triangles), Sonkajärvi Subroute (filled diamonds), Iisalmi Route central basins (empty diamonds), Nilsia
745 Route (filled stars), Kallavesi-Sorsavesi area (empty stars), and Rautalampi Route (black circles) and **b**) diatom taxa
746 (gray circles) and Quaternary deposits. The percentage of fields in a subcatchment (dashed arrow) is plotted passively
747 on top of the ordinations. See Online Resource 2 for further information on the sampling sites and the subcatchment
748 Quaternary deposits

749

750 **Fig. 5 a)** Temporal changes in the diatom-inferred total phosphorus (DI-TP) concentrations of Lake Kirmanjärvi, DCA
751 sample scores of the first and second ordination axes, and the minimum squared chord distances (SCD) between the
752 fossil and modern samples. SCD values below the 2.5 % dashed line indicate that the fossil samples have very good
753 modern analogs in the spatial surface sediment data set, whereas samples between the 2.5 % and 5 % dashed lines have
754 good modern analogs. Samples over the 10 % dashed line are not considered to have modern analogs in the spatial data.
755 The local diatom assemblage zones (LDZ), determined by Kauppila et al. (2012), refer to notable changes in the diatom
756 community of the lake. **b)** Comparison between the reconstructed DI-TP (squares), monitored TP (solid line), and the
757 LOESS smoothed annual averages of the monitored TP values (dashed line)

758

759 **Fig. 6** Fossil Lake Kirmanjärvi samples (black circles) plotted passively on top of a CCA ordination of modern samples
760 (grey circles), water quality and basin variables (arrows). The small diagram on the bottom right shows the trajectory of
761 the fossil samples from the mid-16th century (square) to 2006 (triangle)

762

763 **Electronic supplementary material captions**

764

765 **Online Resource 1** Lake water quality and coordinates at the sampling sites together with basin characteristics

766

767 **Online Resource 2** Quaternary deposits in the subcatchments of the study lakes

768

769 **Online Resource 3** Predicted versus observed values and a residual plot for the total phosphorus (TP) transfer function